Adaptive finite element simulation of stack pollutant emissions over complex terrains

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ABSTRACT

A three-dimensional finite element model for the pollutant dispersion is presented. In these environmental processes over a complex terrain, a mesh generator capable of adapting itself to the topographic characteristics is essential. The first stage of the model consists on the construction of an adaptive tetrahedral mesh of a rectangular region bounded in its lower part by the terrain and in its upper part by a horizontal plane. Once the mesh is constructed, an adaptive local refinement of tetrahedra is used in order to capture the plume rise. Wind measurements are used to compute an interpolated wind field, that is modified by using a mass-consistent model and perturbing its vertical component to introduce the plume rise effect. Then, we use an Eulerian convection–diffusion–reaction model to simulate the pollutant dispersion. In this work, the transport of pollutants is considered and dry deposition is formulated as a boundary condition. The discretization of the stack geometry allows to define the emissions as boundary conditions. The proposed model uses an adaptive finite element space discretization, a Crank-Nicolson time scheme, and a splitting operator. This approach has been applied in La Palma island. Finally, numerical results and conclusions are presented.

1. Introduction

Numerical simulation of pollutant transport and reaction on atmosphere has been the result of important advances in the last thirty years. However, nowadays it remains as a scientific challenge. Key analysis is related to acid rain, ozone, particle matter and toxic emissions [1]. Air quality modelling systems mainly involve three components: emissions, meteorology and transport–chemistry. The first component characterizes different emission sources of chemical compounds [2], the second one determines the atmospheric phenomena as wind and temperature fields, and the third one simulates the transport of pollutants (convection and diffusion) and their chemical reactions.

Environmental Protection Agency (www.epa.gov) classifies Air Quality Modelling Systems as Dispersion, Photochemical and Receptor models. Dispersion models estimate pollutant concentrations at ground level near a punctual source. Photochemical models consider all the sources in a large area. Receptor models identify and characterize the emission sources using receptor measures.

Main dispersion models include Gaussian plume models [3], particle tracking models [4], and puff models [5–7]. The two latter models are based on a Lagrangian approach and usually include first order chemistry reactions or linearised photochemical-reaction models. Versions with complex non-linear photochemical-reaction models have been also developed [8].

In contrast with dispersion models, photochemical ones follow an Eulerian description of the coupled transport and hydrodynamic problem. Most of these Eulerian models approximate the solution to this problem by using finite difference schemes. The modelling domains usually vary from about few thousands to tens of kilometres, and the grid spatial discretization varies from about tens of kilometres to 1 km using nested sub-grids [9–11]. The height of the computational domain usually varies from 4 up to 10 km with a non-uniform vertical discretization between 10 and 20 levels.

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Vertical spatial resolution ranges from few tens metres close to the ground level to one thousand metres above 2 km over the terrain. The number of grid points in this kind of problems can vary from tens to hundreds thousand. The most advanced photochemical models consider local emissions and are known as Plume in Grid (PIG). For this purpose, CAMx [12] and CMAQ [13] include a puff model, and UAM-V [14] uses a plume model.

In environmental management, dispersion and photochemical models have very different applications. The first ones are usually applied to local emission impact assessments, and the second ones to regional planning and monitoring. Both models have a clearly different application scale and photochemical reaction complexity. However, because of the great coupling among key components as ozone, nitrogen compounds, and Volatile Organic Components [15,16], and due to the awareness about the socio-economic impacts of their immission [17–19], some references about the need of coupling local emissions using regional planning with Plume in Grid photochemical models can be found [9,10,20]. Although Plume in Grid models can couple local scales (up to a resolution of 1 km) with regional scales, several limitations have been reported [9]. For example, it is clear that this local high resolution can be insufficient for complex terrains. Thus, the search of alternatives is justified.

In this paper, we present a new methodology for local scale air quality simulations by using a non-steady finite element method with unstructured and adaptive tetrahedral meshes. The aim of this proposal is to introduce an alternative to the standard implementation of current models, improving the computational cost of methods that use structured meshes [21].

Three remarkable uses of unstructured meshes in atmospheric pollution problems are the two-dimensional regional–global examples presented in Lagzi et al., [22], Ahmad et al. [23], the three-dimensional regional examples, including local refinement with element sizes of 2 km, presented in Tomlin et al. [24], and the three-dimensional tetrahedral meshes for local wind field analysis with element sizes ranging from 2 m up to 2 km, see Montenegro et al. [25], Montero et al. [26]. The ideas of this last approach are considered to determine the wind field, that includes the effect of the plume rise, used in the air quality problem.

The wind field is crucial for the pollutant transport, specially in complex terrain areas. In order to simulate it, we have used a mass-consistent model. Several two-dimensional [27] and three-dimensional [28–30] adaptive finite element solutions have been developed by the authors.

The convection, diffusion and reaction problem is usually solved using splitting schemes [31,32] and specific numerical solvers for time integration of photochemical reaction terms [23,33,34]. A non-steady and non-linear transport model is presented in this paper. A stabilized finite element formulation, specifically Least-Squares, with a Crank-Nicolson temporal integration is proposed to solve the problem [35,36]. The chemistry is simulated by using the RIVAD/ARM3. This is a very simplistic empirical chemistry model used in codes such as CALPUFF [37] and is employed here as a proof of concept. More complete chemistry models will be included in future work. The transport and chemical terms are treated separately with Strang splitting operators [38]. The non-linear chemical part is solved node by node with a second order Rosenbrock method [39]. A previous description of the proposed procedure can be found in Pérez-Foguet et al. [40].

In the second Section of the paper we describe the proposed methodology. It mainly involves the following steps: an automatic tetrahedral mesh generator, the wind field simulation, the plume rise approximation and the air pollution simulation. In the third Section we apply the proposed air quality model in La Palma island (Canary Islands, Spain).

2. Wind and air pollution modelling

In this section we introduce the models to simulate the wind field and the transport and reaction of pollutants. The evaluation of the wind field is based on a mass-consistent model, while the transport of pollutant is calculated by a convection–diffusion–reaction PDE by using a non-linear chemical model. The proposed methodology is summarized in Algorithm 1.

Algorithm 1. Wind and air pollution modelling.

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- 1: Construct an adaptive tetrahedral mesh of the domain
 - 1.1: Adaptive discretization of the terrain surface
 - 1.2: Vertical spacing function and 3D distribution of points
 - 1.3: Three-dimensional mesh generation
 - 1.4: Mesh optimisation
 - 2: Wind field simulation from experimental or forecasting data
 - 2.1: Construction of the initial interpolated wind field
 - 2.2: Approximation of the wind field with a mass-consistent model
 - 3: Wind field modification including the plume rise effect
 - 3.1: Compute a plume rise trajectory
 - 3.2: Mesh refinement along the plume rise trajectory
 - 3.3: Simulate a wind field in the new mesh (applying step 2)
 - 3.4: Modify vertical components of the wind field along the plume rise
 - 4: Air pollution simulation from stack emission data
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The paper is organised as follows. In Section 2 we describe the main steps of the proposed methodology. Results are shown in Section 3, and finally the conclusions and future work are presented in Section 4.

2.1. Automatic tetrahedral mesh generation

The studied domain is limited at the bottom by the terrain and at the top by a horizontal plane. The lateral walls are formed by four vertical planes. A uniform distribution of nodes is defined on the upper boundary. A refinement/derefinement algorithm [41] is applied on this uniform mesh to construct a node distribution adapted to the terrain surface and stacks. Once the node distribution is defined both on the terrain and the upper boundary, we distribute the nodes located between both layers by using a vertical spacing function. Next, a three-dimensional mesh generator based on Delaunay triangulation [42] is applied. Finally, the untangling and smoothing procedure described in [43] is used to get a valid mesh and to improve its quality. A detailed description of the mesh generation procedure can be seen in [44,45].

2.2. Wind field simulation

Once the tetrahedral mesh is constructed, we consider a mass-consistent model [28–30] to compute a wind field $\mathbf{u}$ in the three-dimensional domain Ω , with a boundary $\Gamma = \Gamma_a \cup \Gamma_b$, that verifies the continuity equation (mass conservation) for constant density and the impermeability condition on the terrain Γ_a ,

$$\begin{aligned} \nabla \cdot \mathbf{u} &= 0 \quad \text{in } \Omega \\ \mathbf{n} \cdot \mathbf{u} &= 0 \quad \text{on } \Gamma_a \end{aligned} \quad (1)$$

where $\mathbf{n}$ is the outward-pointing normal unit vector.

The model formulates a Least-Squares problem in the domain Ω to find a wind field $\mathbf{u} = (u, v, w)$, such that it is adjusted as much as possible to an interpolated wind field $\mathbf{u}_0 = (u_0, v_0, w_0)$. The wind field $\mathbf{u}$ verifies the Euler–Lagrange equation,

$$\mathbf{u} = \mathbf{u}_0 + \mathbf{P}^{-1} \nabla \phi \quad (2)$$

where ϕ is the Lagrange multiplier, $\mathbf{P}$ is a 3×3 diagonal matrix with $P_{1,1} = P_{2,2} = 2\alpha_1^2$ and $P_{3,3} = 2\alpha_2^2$, being α_1 and α_2 constant in Ω .

The variational approach results in an elliptic problem in ϕ , by substituting (2) in (1), that is solved by using the finite element method.

$$-\nabla \cdot (\mathbf{P}^{-1} \nabla \phi) = \nabla \cdot \mathbf{u}_0 \quad \text{in } \Omega \quad (3)$$

$$-\mathbf{n} \cdot \mathbf{P}^{-1} \nabla \phi = \mathbf{n} \cdot \mathbf{u}_0 \quad \text{on } \Gamma_a \quad (4)$$

$$\phi = 0 \quad \text{on } \Gamma_b \quad (5)$$

The interpolated wind field $\mathbf{u}_0$ can be constructed from experimental data or meteorological forecasting models. In this paper we consider the first case. Therefore, we use a horizontal interpolation and a vertical extrapolation of the available measurements to construct $\mathbf{u}_0$ in the whole computational domain. The horizontal interpolation is formulated as a function of the inverse of the squared distance to the measurement stations, and the inverse of the height differences [28]. For the vertical extrapolation, a log-linear wind profile is considered. It takes into account the horizontal interpolation and the effect of roughness on the wind velocity [46–50]. For proof of concept this work adopts the simpler Pasquill stability class. Future work can employ a more precisely planetary boundary layer theory. Note that the log law is recommended only for bi-dimensional fluid flow. In three-dimensional fluid flows there is a significant change on the near wall behaviour. The mean velocity and the shear stress vector may shift directions with the distance from the wall; furthermore, the shear stress and the mean vector are not, generally, aligned [51]. However, we use the log law because there is no information about the near three-dimensional wall behaviour, and it is a simple method to construct a reasonable interpolated wind field that will be modified by the mass-consistent model taking into account the topography.

2.3. Plume rise

The plume rise phenomenon is mainly due to the difference of temperature between the released substance and the environment air, and the initial momentum. The trajectory of the plume rise has been widely studied in the past [52–54]. These works differentiates two kind of cases: *predominant buoyancy rise* and *predominant momentum rise*. The characterization of these types essentially depends on the ratio between the intensities of the pollutant emission velocity and the wind velocity at the top of the stack.

The previous studies are based on an important assumption such as a constant ambient wind field with null vertical component. This assumption produce that the mean trajectory of the plume is placed on a vertical plane, that contains the stack. Therefore, the resulting plume rise trajectory cannot be directly applied in the case of complex terrains. In this case, the plume rise trajectory follows a bent curve due to the variation of the ambient wind field.

The first objective of this Section is to construct a reasonable plume rise mean trajectory that takes into account the Briggs' equations and the ambient wind field obtained in the previous Section. In the future work we will implement a more precise numerically integrated approach, as the one used in the PRIME algorithm [55].

Once the plume trajectory is defined, the effect of the pollutant emission will be introduced by modifying the ambient vertical wind velocity (w) along the region of the plume rise trajectory, taking into account the vertical velocity of the plume (w_0). As stated by Briggs [53] w_0 is a convenient variable with which to identify the plume. In this Section we introduce a method to give an

approximation of w_0 at any point in the surroundings of the bent plume trajectory.

2.3.1. Plume rise mean trajectory

Using the Briggs formulation, we know the predominant force in the plume rise, and the values of the effective height of the plume (z_H) and the horizontal distance (d_f). In this section a formulation to know the trajectory of the plume rise and the vertical component of the perturbed wind field along this trajectory has to be defined.

Firstly, we will consider separately the two different types of trajectories, bent-over or vertical, in the vertical plane defined by the top of the stack and the end of the plume.

Secondly, we will extend the previous results to three-dimensional trajectories in the case of a complex terrain.

2.3.1.1. Bent-over plume mean trajectory in a vertical plane. In all cases with d_f different from zero, the driving force is buoyancy, except for stable conditions and calm wind. In order to know the plume rise trajectory, we propose to combine a horizontal and a vertical motion, verifying certain known conditions.

The vertical motion along the mean trajectory of the plume is defined by an acceleration $a_0(t)$, a velocity $w_0(t)$ and $z(t)$, from the initial time $t = 0$ to the final time $t = t_f$ when the plume reaches the effective height, verifying the following conditions

$$z(0) = z'_c \quad z(t_f) = z_H \quad (6)$$

$$w_0(0) = w_c \quad w_0(t_f) = 0 \quad (7)$$

Since there are four conditions on the vertical motion, we propose a cubic approximation of $z(t)$, and therefore a quadratic approximation of $w_0(t)$, and a linear approximation of $a_0(t)$ given by the following expressions

$$z(t) = z'_c + w_c t + \frac{-2w_c t_f + 3(z_H - z'_c)}{t_f^2} t^2 + \frac{w_c t_f - 2(z_H - z'_c)}{t_f^3} t^3 \quad (8)$$

$$w_0(t) = w_c + \frac{-4w_c t_f + 6(z_H - z'_c)}{t_f^2} t + \frac{3w_c t_f - 6(z_H - z'_c)}{t_f^3} t^2 \quad (9)$$

$$a_0(t) = \frac{-4w_c t_f + 6(z_H - z'_c)}{t_f^2} + \frac{6w_c t_f - 12(z_H - z'_c)}{t_f^3} t \quad (10)$$

To satisfy the conditions (7), and considering that $w_0(t)$ is a decreasing function in $[0, t_f]$, we impose $a_0(t) \leq 0$ along the trajectory, obtaining the following condition for the unknown t_f

$$\frac{3}{2}(z_H - z'_c) \leq w_c t_f \leq 3(z_H - z'_c) \quad (11)$$

To determine the value of t_f , an additional assumption has to be done introducing the horizontal motion.

The horizontal motion is defined by a uniformly accelerated motion, with a constant positive acceleration vector $\mathbf{a}_d = (a_{dx}, a_{dy})$, a velocity $\mathbf{u}_d(t) = (u_d(t), v_d(t))$, and a horizontal relative position vector $\mathbf{d}(t) = (x(t) - x_c, y(t) - y_c)$ with respect to the centre of the stack, verifying the following conditions

$$|\mathbf{d}(t_f)| = d_f \quad (12)$$

$$\mathbf{u}_d(0) = \mathbf{u}(x_c, y_c, z_c) \quad (13)$$

Note that at the top of the stack the ambient wind $\mathbf{u}(x_c, y_c, z_c)$ has only horizontal components because of the imposition of the impermeability condition (1) at the top of the stack when solving the mass-consistent problem.

The horizontal trajectory can be expressed as

$$\begin{aligned} x(t) &= x_c + u(x_c, y_c, z_c) t + \frac{1}{2} a_{dx} t^2 \\ y(t) &= y_c + v(x_c, y_c, z_c) t + \frac{1}{2} a_{dy} t^2 \end{aligned} \quad (14)$$

From the previous conditions, we obtain an additional restriction for the unknown t_f

$$t_f = \frac{1}{a_d} \left(-|\mathbf{u}(x_c, y_c, z_c)| + \sqrt{|\mathbf{u}(x_c, y_c, z_c)|^2 + 2a_d d_f} \right) \quad (15)$$

where $a_d = |\mathbf{a}_d|$ is still an unknown. Imposing the condition (11) we obtain the following expression for a_d

$$a_d = (1 + \delta) \frac{2w_c}{3(z_H - z'_c)} \left[(1 + \delta) \frac{w_c}{3(z_H - z'_c)} d_f - |\mathbf{u}(x_c, y_c, z_c)| \right] \quad (16)$$

being $0 \leq \delta \leq 1$. For $\delta = 0$, the value of t_f is related to the upper bound in (11) and, for $\delta = 1$, to the lower bound. The case $\delta = 1/2$ corresponds to a value of t_f which produces a constant vertical acceleration a_0 , a linear vertical component of velocity $w_0(t)$ and a quadratic vertical position $z(t)$.

Note that if the parameter δ is fixed, the horizontal and vertical motion is completely defined verifying strictly the conditions (12), (13), (6) and (7).

In this paper we will consider the value $\delta = 1/2$ due to the considerations presented at the end of the following Section.

2.3.1.2. Vertical plume trajectory. In all cases where d_f is equal to zero, the horizontal motion of the plume until reaching the effective height can be considered negligible. Thus the trajectory of the gases is nearly vertical.

In this case, we propose a vertical motion along the trajectory of the plume with a constant negative acceleration a_0 , a linear velocity $w_0(t)$ and a quadratic trajectory $z(t)$. Imposing the conditions (6) and (7) this vertical motion is completely defined.

$$t_f = \frac{2(z_H - z'_c)}{w_c} \quad (17)$$

$$a_0 = \frac{-w_c}{t_f} \quad (18)$$

$$w_0(z) = w_c \sqrt{1 - \frac{2(z - z'_c)}{w_c t_f}} \quad (19)$$

We remark that using the value of t_f from (17) in the Equation (8), the coefficient of its cubic term is null. As we have commented in the previous Section this cancellation also occurs when $\delta = 1/2$ is chosen in (16). It is also important to note that using the value of t_f from (17) in the equation (10) results (18). Thus, fixing $\delta = 1/2$, the vertical motion in the bent-over plume case, given by Equations (8)–(10), and in the vertical plume case, given by Equations (17)–(19), are the same. This result makes compatible the vertical motion in both cases and justifies the election of $\delta = 1/2$.

2.3.1.3. Extension of the plume mean trajectory to 3D over complex terrains. Several approximations can be considered to define the mean trajectory of bent curved plumes considering influence of complex terrains. Extension of vertical plumes is straightforward. In this paper, we propose a new solution that takes into account the trajectories obtained in the vertical plane case, and the non uniform three-dimensional ambient wind obtained in Section 2.2.

A three-dimensional bent trajectory by an iterative process starting from the emission point (x_c, y_c, z_c) . We uniformly subdivide the interval of time $[0, t_f]$ defining $n+1$ time instants $t_0, t_1, \dots, t_n$, such that $t_i = i\Delta t$ for $i = 0, 1, \dots, n$. For each time instant t_i a position of the mean trajectory is calculated.

Starting in time t_0 at position $\mathbf{x}_0 = (x_c, y_c, z_c)$, and known the point $\mathbf{x}_i = (x_i, y_i, z_i)$, $\mathbf{x}_{i+1}$ can be computed from $\mathbf{x}_i$ with the following formulae

$$\mathbf{x}_{i+1} = \mathbf{x}_i + \frac{\mathbf{u}(x_i, y_i, z_i)}{|\mathbf{u}(x_i, y_i, z_i), \mathbf{v}(x_i, y_i, z_i)|} \left(\mathbf{u}(x_c, y_c, z_c) \Delta t + \frac{1}{2} a_d \Delta t^2 \right) \quad (20)$$

$$y_{i+1} = y_i + \frac{v(x_i, y_i, z_i)}{|\mathbf{u}(x_i, y_i, z_i), \mathbf{v}(x_i, y_i, z_i)|} \left(v(x_c, y_c, z_c) \Delta t + \frac{1}{2} a_d \Delta t^2 \right) \quad (21)$$

$$z_{i+1} = z(t_i) \quad (22)$$

where $z(t_i)$ is defined by (8), and the horizontal movement is deduced from (14).

The mean trajectory obtained by this method is a three-dimensional polygonal line, such that the longitude of its projection on the horizontal plane approximates the longitude d_f . In addition, the final height z_n coincides with the effective height z_H . Therefore, this method tries to verify the main values of the end of the plume considering Briggs' equations.

Fig. 1. La Palma tetrahedral mesh before (a) and after (b) the refinement under 3000 m.

Fig. 2. Mesh detail around the stack before (a) and after (b) the refinement.

Fig. 3. Mesh quality before (a) and after (b) the refinement.

2.3.2. Local mesh refinement along a Gaussian plume

In order to modify the ambient vertical wind velocity (w) along the region of the plume rise, we need to have a sufficient mesh resolution in this area. For this reason, we propose to refine locally the mesh along the Gaussian plume [56] until all the tetrahedra inside that region verify a size criteria.

The region can be defined in function of the diffusion coefficients of the Gaussian plume (σ_y, σ_z). To decide if a sample point $\mathbf{x}$ of a tetrahedron T is inside this region, the first step is to find its closest point $\mathbf{x}_p \in \mathbf{x}_0, \mathbf{x}_1, \dots, \mathbf{x}_n$. The second step is to calculate the distance $d(t_p) = \sqrt{(x(t_p) - x_c)^2 + (y(t_p) - y_c)^2}$ where the time $t_p = p\Delta t$, and $x(t_p), y(t_p)$ are evaluated with Equation (14). Then, we can compute the values $\sigma_z(d(t_p))$ and $\sigma_y(d(t_p))$ in function of the Pasquill stability class. These values define vertical ellipses (centred in $\mathbf{x}_p$) whose major and minor semi-axes are proportional to σ_y and σ_z respectively. Finally, fixing a value $K_1 > 1$, we consider that the sample point $\mathbf{x}$ of the tetrahedron T is inside the Gaussian plume region if $|\mathbf{x} - \mathbf{x}_p| < K_1 \sigma_y$.

For the refinement of the marked tetrahedra, we use a local refinement algorithm [57] based on the 8-subtetrahedron subdivision developed in [58].

2.3.3. Wind field perturbation along the plume rise

Finally, a new ambient wind field $\mathbf{u}$ is obtained on the refined mesh with the mass-consistent model described in Section 2.2. The

Fig. 4. Location of the four measurement stations and its corresponding wind velocity. The colourmap represents the wind velocity in (m s^{-1}).

Fig. 5. Interpolated (a) and resulting (b) wind field at 20 m over the terrain. The colourmap represents the wind velocity in (m s^{-1}). (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

effect of the gas emission is introduced in this field by modifying its vertical component.

In the case of bent curve trajectory, the vertical component of the ambient wind velocity $\mathbf{u}$ is modified at any point $\mathbf{x}$ of the domain Ω located inside the region that was defined in the previous Section. More precisely, the vertical component w of the velocity $\mathbf{u}$ in x is imposed by using Equation (9) as $w_0(t_p)$, being x_p the closest point to $\mathbf{x}$ in the discrete trajectory $\mathbf{x}_0, \mathbf{x}_1, \dots, \mathbf{x}_n$ of the plume.

In the case of vertical trajectory, the vertical component of the ambient wind velocity is modified inside a standard cylinder

Table 1
Location in UTM, and measurements of the wind velocity at the four stations.

Station	X	Y	$ V $ (m s^{-1})	$\varphi(V)$ ($^\circ$)
LPA	213,000	3,172,000	3	225
MBI	224,760	3,180,000	3	270
MBII	226,500	3,175,000	3.5	265
MBIII	222,000	3,175,500	3	245

defined in the previous Section. That is, given a point $\mathbf{x} = (x, y, z)$ inside the cylinder, we impose the vertical velocity w of $\mathbf{u}$ by using Equation (19) as $w_0(z)$.

2.4. Air pollution simulation

The air pollution simulation consists in solving the unsteady convection–diffusion–reaction formulation with a stabilized finite element method, specifically Least-Squares method, with a Crank–Nicolson temporal discretization. The equation system governing the problem can be expressed with the following vectorial equation:

$$\frac{\partial c}{\partial t} + \mathbf{u} \cdot \nabla c = \nabla \cdot (K \nabla c) + e + s(c) \quad (23)$$

for the spatial coordinates x and time t , $(x, t) \in \Omega \times (0, t^{\text{end}}]$, with initial condition $c(x, 0) = c^{\text{ini}}(x)$ on $x \in \Omega$, and the following boundary conditions:

Fig. 6. Interpolated (a) and resulting (b) wind field at 100 m over the terrain. The colourmap represents the wind velocity in (m s^{-1}). (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

Fig. 7. Interpolated (a) and resulting (b) wind field at 500 m over the terrain. The colourmap represents the wind velocity in (m s^{-1}). (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

Fig. 9. Perturbed wind field streamlines from the top of the stack. The colourmap represents the wind velocity in (m s^{-1}). (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

$$\begin{cases} c(x, t) = c^{\text{emi}}(x) & \text{in } \Gamma_S : \text{Top of the stack} \\ n \cdot K \nabla c = -V^d c & \text{in } \Gamma_R : \text{Terrain} \\ n \cdot \nabla c = 0 & \text{in } \Gamma_{W_{\text{out}}} : \text{Outlet wind boundaries} \\ c(x, t) = c^{\text{out}}(x) & \text{in } \Gamma_{W_{\text{in}}} : \text{Inlet wind boundaries} \end{cases} \quad (24)$$

where ∇ is the gradient with respect x , and c , u , e and $s(c)$ are respectively the concentration, the perturbed wind velocity, the emission and the chemical vectors with a dimension n_c (the number of pollutant species), K is the diffusion matrix of dimension $3 \times n_c$, V^d is the deposition diagonal matrix with dimension n_c , and n is the outward-pointing normal unit vector, c^{emi} is the concentration of the emission in the top of the stack, and c^{out} the outside concentration at the inlet wind boundaries. Scalar product “ $\cdot$ ” is applied n_c times: the first argument is multiplied by each one of the n_c components of the second argument.

The complete description of photochemical reaction of atmospheric species is highly complex [59–62]. For instance, detailed Volatile Organic Components decomposition involves hundreds of thousand reactions [63,64] that need special methodologies to reduce the number of the modelled reactions and species. Reference models for gaseous phase reactions involve some tens of compounds [65,66]. The most simplified models just involve about ten reactive species [67]. On the other side, depending on the

Fig. 8. Refinement steps along the plume rise: (a) 1 step, (b) 6 steps.

application, it can be necessary to take into account aqueous phase reactions, that involve several other reactions and species. The RIVAD/ARM3 model is one of the most simplified models that permit to simulate both processes, aqueous and gaseous, involving transport and reaction of four species [37]. In this paper, we have considered the RIVAD/ARM3 model for the chemical term $s(c)$.

2.4.1. Linear chemical problem

In this case the chemical term is linear, that is, $s(c) = Ac$ where A is constant matrix. The resulting Equation (23) is solved with a Crank-Nicolson time integration scheme, and a spatial discretization with a stabilized finite element method, Least-Squares. Concentrations c^n and c^{n+1} at times t^n and $t^{n+1} = t^n + \Delta t$ are related using a Crank-Nicolson scheme as $c^{n+1} = c^n + \Delta t/2 [\partial c^{n+1}/\partial t + \partial c^n/\partial t]$. And using Least-Squares we obtain a symmetric problem that can be written as the following equation system

$$Bc^{n+1} = f \quad (25)$$

where c^{n+1} is the concentration vector approximation at t^{n+1} in the degrees of freedom of the finite element discretization, f depends on c^n and pollutant emissions, and B is a square matrix with dimension $(n_c \times n_{dof})$, being n_{dof} the number of degrees of freedom.

In order to solve this linear system it is necessary to find an efficient solver, using a sparse matrix storage. Since B is a symmetric positive definite matrix, we have considered a solver based on a conjugate gradient method preconditioned with an incomplete Cholesky factorisation density type [68]. More details about the implementation of this system equation solver can be found in [35] and [36].

2.4.2. Non-linear chemical problem

To deal with the non-linearity of the reactive term in the convection–diffusion–reaction Equation (23), we have considered a splitting method that separates this equation into a convection–diffusion equation and a reaction equation. We will make use of the second order splitting operator (*Strang splitting*) proposed by [38]. Using this method the complete equation is splitted into two equations. The convection–diffusion equation that is solved using

Fig. 10. Evolution of the immission concentration distribution (g m^{-3}) of the primary pollutant SO_2 after (a) 5 min, (b) 10 min, (c) 15 min, (d) 20 min, (e) 25 min and (f) 30 min.

the same method proposed in the previous Section, being $A = 0$. And the non-linear chemical equations that are solved node by node with a second order Rosenbrock method (ROS2) [39]. To use the ROS2 method, the Jacobian square matrix of $s(c)$ of dimension n_c has to be computed.

3. Results

In this section a realistic example is presented by using the methodology described in this paper. We propose to study a region in La Palma island, where several wind measurements data are given, and SO_2 and NO_2 emissions from a stack are considered. The topography of the island is real, from a digital elevation model, but the wind field measurements, and the stack location and emissions, are simulated.

3.1. Adaptive tetrahedral mesh

The studied domain taken under consideration is a rectangular area with dimensions $15,600 \times 22,803$ m. The topography of the

terrain is highly complex ranging from the sea level up to a maximum height of 2279 m with several deep valleys. The upper boundary of the domain has been placed at $h = 9000$ m. The digital elevation model of the area is defined over a uniform grid with a spacing step of 200 m in directions x and y . We add the stack geometry to the topographical data. Let us consider a stack with a height of 150 m over the terrain and the diameter at its top of 15 m, the location of the stack is UTM 28N 224,760 3,178,743, and is shown in Fig. 4. The mesh must be able to detect the details of the stack. Therefore, we impose an element size of about 2×2 m in the stack. If we start from a uniform mesh τ_1 of the rectangular area with an element size of about 2000×2000 m, we must make ten refinement steps close to the stack. However, we make only five global refinement steps over the whole domain and five additional local refinements in the surroundings of the stack. We have considered the derefinement parameter of $\epsilon = 40$ m. Thus, the adapted mesh nears the terrain surface with an error less than that value. On the other hand, the node distribution of τ_1 is the one considered on the upper boundary of the domain.

Fig. 11. Evolution of the immission concentration distribution (gm^{-3}) of the secondary pollutant SO_4 after (a) 5 min, (b) 10 min, (c) 15 min, (d) 20 min, (e) 25 min and (f) 30 min.

Fig. 1 (a) shows the resulting tetrahedral mesh, with 198,146 nodes and 1,123,012 tetrahedra. Fig. 2 (a) shows a detail of the surroundings of the stack. Fig. 3 (a) shows the distribution of the mesh element quality using the mean ratio algebraic shape quality metric.

In order to have a more suitable mesh for the air quality problem, all the elements located below the height 3000 m have been refined. Fig. 1 (b) shows the refined mesh with 455,953 nodes and 2,622,454 tetrahedra. Fig. 2 (b) shows a detail of the surroundings of the stack and Fig. 3 (b) shows the distribution of the mesh element quality.

3.2. Ambient wind field

The wind field is constructed from the measurements of 4 stations. The locations of these stations are shown in Fig. 4, and the values of the corresponding wind velocities are listed in Table 1. In the application of the mass-consistent model, described in Section 2.2, we have considered a neutral Pasquill stability class, and the following parameter values: $u_g = (-10, -1, 0)$, $\alpha = 1.0$, $\xi = 0.5$, $\gamma = 0.15$.

Figs. 5–7 show the interpolated (u_0) and the resulting (u) ambient wind field at heights 20 m, 100 m and 500 m over the

terrain. The interpolated wind field preserves the measurements of the stations, but the terrain has no impact over it. Therefore, it can be seen in the figures how the wind field u_0 crosses the terrain. However, the terrain has a strongly effect on the resulting wind field, verifying the conditions of the mass-consistent model. It can be noted that the wind field velocity is higher in the ridges of the mountains. In addition, the wind field is channelled in the ravines. It also can be noted, that the effect of the terrain decreases as the elevation increases, and the value of the wind velocity increases and tends to the uniform geostrophic wind field.

3.3. Plume rise

The stack emission characteristics are given by a gas velocity of $w_c = 5 \text{ m s}^{-1}$ at a temperature of $T_c = 573 \text{ K}$, the ambient temperature is $T = 298 \text{ K}$, and the internal exit diameter is $D_c = 15 \text{ m}$.

The velocity of the ambient wind at the top of the stack ($z_c = 1454 \text{ m}$) has a magnitude of $|u_c| = 7.13 \text{ m s}^{-1}$, with horizontal components $u_c = -7.09 \text{ m s}^{-1}$ and $v_c = -0.77 \text{ m s}^{-1}$. So, a predominant buoyancy rise with a bent plume trajectory results in this case.

Fig. 12. Evolution of the isosurface corresponding to a concentration of $10 \mu\text{g m}^{-3}$ for the primary pollutant SO_2 after (a) 5 min, (b) 10 min, (c) 15 min, (d) 20 min, (e) 25 min and (f) 30 min.

With the previous values, using the Briggs' equations, we obtain $z'_c = 1406$ m, $F = 5390$ m⁴ s⁻³, $z_H = 2347$ m and $d_f = 3700$ m.

The mean trajectory of the plume is then determined by using the method proposed in Section 2.3.1. Fig. 8 shows the effect of the local refinement along the Gaussian plume after one (a) and six (b) refinement steps. The final mesh resolution along the plume is enough to capture the transport of pollutants. Note that the plume follows a bent trajectory.

Once the mesh has been refined, a new mass-consistent wind field is computed in the new mesh. In order to capture the buoyancy rise effect, the resulting ambient wind field is perturbed modifying its vertical component as proposed in Section 2.3.3. Fig. 9 represents the streamlines that start at the top of the stack. These streamlines follow a reasonable bent curve trajectory, that introduces the main physical effects of the emission problem over a complex terrain.

3.4. Air pollution

At this point we can compute the air pollution in the studied region. For the non-linear chemical part, the RIVAD/ARM3 model has

been used. We simulate the concentrations of four species $c_1 = [\text{SO}_2]$, $c_2 = [\text{SO}_4]$, $c_3 = [\text{NO}_2]$ and $c_4 = [\text{NO}_3]$. The stack emission concentration (c^{emi}) for the primary pollutants is fixed to $c_1^{\text{emi}} = c_3^{\text{emi}} = 6$ g m⁻³, and for the secondary pollutants $c_2^{\text{emi}} = c_4^{\text{emi}} = 0$. We have considered a horizontal diffusion coefficient of 8×10^{-6} m² s⁻¹ and a vertical diffusion coefficient of 4×10^{-6} m² s⁻¹ for the four species, corresponding to a mean value for the air. All the four diagonal terms of the deposition matrix (V^d) have been fixed to 1.3×10^{-3} m s⁻¹. In addition, we have fixed null values for vectors e , c^{ini} and c^{out} . The time-dependent problem has been simulated with a time step $\Delta t = 10$ s.

Figs. 10 and 11 show the evolution of the immission concentration distributions for SO₂ and SO₄, respectively. These Figures are interesting since they give information about the pollutant concentration at the ground-level. Note that primary pollutant tends to have the highest concentrations near the emission source, while the highest concentrations of the secondary pollutant are located further. This is consistent with the chemical reaction effect.

Figs. 12 and 13 show the evolution of the isosurfaces corresponding to a concentration value of 10 µg m⁻³ for SO₂ and SO₄,

Fig. 13. Evolution of the isosurface corresponding to a concentration of 10 µg m⁻³ for the secondary pollutant SO₄ after (a) 5 min, (b) 10 min, (c) 15 min, (d) 20 min, (e) 25 min and (f) 30 min.

Fig. 14. Wind field at 10 m over the terrain, and location of the stack and the points A–E.

respectively. Note that the secondary pollutant is more prevalent than the primary one in the front of the plume at the final time steps. It also can be noted that the main effect of the plume dispersion includes the orography, the wind and the emissions.

The evolution of the immission concentration has been studied in five points (A–E). Fig. 14 shows their locations. These points are grouped in two separate sets. The horizontal distances from the stack to the first two points (A,B) are approximately 100 m and 1000 m, respectively. We note that there are no obstacles between the stack and this first group of points. The second group is formed by three points (C,D,E) that are located at different heights behind the crest of a mountain. The approximated horizontal distance from the stack to this three points are 3000 m, 4000 m and 5000 m respectively.

Fig. 15 shows the evolution of the immission concentration at the points (A–E). It can be observed the influence of the distance to the stack and the orography. The numerical results verify the qualitative standard behaviour in the first group of points (concentration decreases with the distance). Moreover, reasonable values are reached behind the crest of the mountain. Note that the value of the steady concentration at the deepest point (C) is lower than the values at points D and E.

A final comment about the computational complexity of the evolution process should be done. For each time step we have to solve a finite element problem with a number of degrees of freedom about the number of nodes multiplied by the number of species, i.e. $455,953 \times 4 = 1,823,812$. The number of time steps in

the simulation period (about 30 min) is $30 \times 60/10 = 180$. Therefore, in the whole evolution process about 180 linear equation systems with 1,823,812 unknowns must be solved. The computational cost corresponding to the mesh generation, wind simulation, and the resolution of the ordinary differential equations in the splitting method are insignificant with respect to the resolution of the unsteady convection–diffusion equation. In a computer with 128 GB of RAM memory and 2.34 GHz, the total computing time is about 40 min. In a future work, the present computational complexity will be significantly reduced by using a refinement/derefinement strategy that follows the front of the pollutant plume, minimising the number of degrees of freedom in each time step.

4. Conclusions

We have introduced a new methodology for solving air pollution problems over a complex terrain. The adaptive three-dimensional mesh generation discretize domains defined over complex terrains which include several stack geometries, with a minimal user intervention and low computational cost. The mass-consistent model obtains an ambient wind field that takes into account the complex orography. The local mesh refinement along the Gaussian plume, allows to perturb the ambient wind field to introduce the effect of the pollutant emissions. The convection–diffusion–reaction equation obtains the values of concentration for all the pollutants in the whole three-dimensional domain. The proposed methodology uses the necessary mesh resolution to consider the main effects of the physical phenomena that are involved in air pollution process. The numerical experiment shows a reasonable behaviour of the proposed method. The present manuscript introduces a solid framework to develop more accurate models in the future.

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Fig. 15. Immission concentration evolution of the primary pollutant at points A–E.

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